

DESCRIPTION OF MAP UNITS

Major lithology

QTf **YOUNGER BASIN-FILL DEPOSITS** (Holocene to Pliocene)—Unconsolidated to semi-consolidated, poorly sorted to well-sorted, coarse- to fine-grained deposits of alluvial fans, flood plains, lowlands, playas, and Pleistocene lakes

Tf **OLDER BASIN-FILL DEPOSITS** (Pliocene and upper Miocene)—Conglomerate, sandstone, siltstone, mudstone, freshwater limestone, evaporite beds, tuff, and interbedded lava flows

Areal extent

QTf and , in many areas, Tf represent erosion products from present-day mountain ranges. QTf underlies present-day valleys of the Great Basin. Alluvial fans and pediments usually underlain by coarse-grained deposits, and fine-grained deposits underlie lowlands. Flood plains are underlain by alternating intervals of sorted coarse- and fine-grained deposits. Tf is found mainly along valley margins and in mountain blocks. This unit, however, may not as yet have been recognized in many areas (Stewart, 1980, p. 95), and may underlie QTf in many basins

Water-bearing properties

These units yield most of the ground water used in the Great Basin. QTf may be more productive than Tf because the latter is usually more consolidated. Deposits below flood plains usually are more productive than those below alluvial fans, which are, in turn, more productive than deposits below lowlands. Aquifer-test data for basins in central and eastern Nevada and western Utah, gathered for the MX missile-siting investigation, were compiled by Bunch and Harrill (1984, p. 115-118). Estimated hydraulic conductivities computed from these test results (transmissivity divided by the length of screened interval in the well) range from 0.02 ft/d (feet per day) in Sever Desert, Utah to 140 ft/d in Railroad and Hot Creek Valleys, Nev. (Bunch and Harrill, 1984, p. 115-118). The average value determined from 18 tests in 14 basins is 74 ft/d. The maximum and minimum values appear to represent a reasonable range of conductivities for basin-fill deposits. The average value, however, may be biased toward higher values because the purpose of the investigation was to find and eventually develop high-yield sources of ground water

QTV **VOLCANIC ROCKS** (Holocene to Eocene)—Lava flows, flow breccias, shallow intrusive rocks, and ash-fall, air-fall, and water-lain tuffs. Compositions range from basalt to rhyolite

Mountain ranges throughout the Great Basin; also underlie or are interbedded with QTf and Tf in many basins

Water-bearing properties generally not known except for basalt aquifers at Fallon, Nev., and Pavant Valley, Utah, and lava flow and tuff aquifers at Nevada Test Site in southern Nev. Estimates of transmissivity are 4,100-170,000 ft²/d (feet squared per day) at Fallon (Glancy, 1981, p. 7), 24,000-2,700,000 ft²/d in Pavant Valley (Mower, 1965, p. 36-38), and 170-9,100 ft²/d at Nevada Test Site (Winograd and Thordarson, 1975, p. 22-23)

TRI **INTRUSIVE ROCKS** (Miocene to Upper Triassic)—Mostly granodiorite and quartz monzonite

Outcrops of these rocks are found in many mountain ranges. Aeromagnetic data indicate that many intrusive bodies are much more extensive in the subsurface than is indicated by outcrop areas

Shallow fractures may be capable of storing and transmitting small quantities of water; however, unit generally is considered to be a barrier to regional ground-water movement

TRc **MARINE AND CONTINENTAL ROCKS** (Lower Miocene to Middle Triassic)—Unit represents diverse lithologies that include continental fluvial, lacustrine, and aeolian deposits and clastic and carbonate sedimentary rocks and volcanic rocks of marine origin

Mountain ranges, mostly near the eastern and western margins of the Great Basin

Water-bearing properties unknown, although transmissivities could range through several orders of magnitude because of the diversity of lithologies

FCc **MARINE SEDIMENTARY AND VOLCANIC ROCKS** (Lower Triassic to Lower Cambrian)—Unit consists of deep-water clastic rocks, chert, and volcanic rocks

Structurally overlies units of approximately equivalent age along the Roberts Mountains and Golconda thrusts in central Nevada (fig. 2)

Water-bearing properties unknown. Probably does not store or transmit much water

PDC **CLASTIC ROCKS** (Upper Permian to Upper Devonian)—Shale, sandstone, and conglomerate¹

Eastern Great Basin

Water-bearing properties largely unknown. Presumed to act as a local barrier to flow between FMld and D-Cld

FMld **CARBONATE ROCKS** (Upper Triassic to Lower Mississippian)—Limestone, commonly silty or sandy, and lesser amounts of dolomite, shale, siltstone, and sandstone. This unit is generally more clastic than D-Cld. In parts of central Utah, proportion of clastic rocks may exceed that of carbonate rocks

Eastern Great Basin

Forms upper part of carbonate-rock aquifer. Water-bearing properties largely unknown. Estimated to be similar to those of D-Cld at the Nevada Test Site (Winograd and Thordarson, 1975, p. 30-31)

Dcsl **CARBONATE AND CLASTIC ROCKS** (Middle Devonian to Middle Cambrian)—Thin bedded to laminated, sandy to silty limestone, sandstone, siltstone, and shale

Central Nevada

Water-bearing properties largely unknown. May be similar to those for units FCc, FMld, and D-Cld

Dcld **CARBONATE ROCKS** (Middle Devonian to Middle Cambrian)—Limestone and dolomite, with lesser amounts of shale, sandstone, and quartzite

Eastern Great Basin as far west as central Nevada

Forms lower part of carbonate-rock aquifer. Estimates of transmissivity are as follows: from 170-5,200 ft²/d at Nevada Test Site (Winograd and Thordarson, 1975, p. 22); 14,000 ft²/d in White River Valley, Nev. (Plume, 1984, p. 624); 200 ft²/d in Steptoe Valley, Nev., 13,400 ft²/d in Dry Lake Valley, Nev., 400 ft²/d in Garden Valley, Nev., and 40,000 and 250,000 ft²/d in Coyote Spring Valley, Nev. (Bunch and Harrill, 1984, p. 119)

Cp-Cc **CLASTIC ROCKS** (Lower Cambrian and upper Precambrian)—Quartzite, shale, siltstone, sandstone, limestone, and dolomite. Metamorphosed in some areas

Eastern Great Basin

Considered impermeable by Winograd and Thordarson (1975, p. 11). Significant secondary permeability probably does not develop below the older unit of carbonate rocks D-Cld, (so this unit C-Pc c) may form a base for the carbonate-rock aquifer of the eastern Great Basin. Barrier to regional flow where it is in fault contact with the carbonate rock is in fault contact with the carbonate rock Probably not capable of storing or transmitting much water. Barrier to regional flow where it is in fault contact with the carbonate-rock aquifer

pCb **BASEMENT ROCKS** (Precambrian)—Gneiss, schist, gneissic granite, and granite

Outcrops near southern, southeastern, and northeastern boundaries of Great Basin. May underlie eastern Great Basin as far west as central Nevada (Stewart, 1980, p. 9-11)

¹This unit mostly represents erosion products from the Antler and Sonoma highlands and from a continental source farther east (Welsch and Bissell, 1979, p. Y21). These deposits accumulated from Late Devonian through Mississippian time in eastern Nevada and western Utah. Farther west, however, the deposits appear to have accumulated through the Permian. This distinction is important because the unit stratigraphically separates FMld and D-Cld in eastern Nevada and western Utah.